



HAL
open science

The Kufrah paleodrainage system in Libya: A past connection to the Mediterranean Sea?

Philippe Paillou, Stephen Tooth, S. Lopez

► **To cite this version:**

Philippe Paillou, Stephen Tooth, S. Lopez. The Kufrah paleodrainage system in Libya: A past connection to the Mediterranean Sea?. *Comptes Rendus Géoscience*, 2012, 344 (8), pp.406-414. 10.1016/j.crte.2012.07.002 . hal-00833333

HAL Id: hal-00833333

<https://hal.science/hal-00833333v1>

Submitted on 12 Jun 2013

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60
61
62
63
64
65

**The Kufrah Paleodrainage System in Libya:
A Past Connection to the Mediterranean Sea ?**

**Le système paléo-hydrographique de Kufrah en Libye :
Une ancienne connexion avec la mer Méditerranée ?**

Philippe PAILLOU

Univ. Bordeaux, LAB,UMR 5804, F-33270, Floirac, France

Tel: +33 557 776 126 Fax: +33 557 776 110

E-mail: philippe.paillou@obs.u-bordeaux1.fr

Stephen TOOTH

Institute of Geography and Earth Sciences, Aberystwyth University, Ceredigion, UK

Sylvia LOPEZ

Univ. Bordeaux, LAB,UMR 5804, F-33270, Floirac, France

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58

Abstract: Paillou et al. (2009) mapped a 900 km-long paleodrainage system in eastern Libya, the Kufrah River, that could have linked the southern Kufrah Basin to the Mediterranean coast through the Sirt Basin, possibly as long ago as the middle Miocene. We study here the potential connection between the terminal part of the Kufrah River and the Mediterranean Sea through the Wadi Sahabi paleochannel, which may have constituted the northern extension of the lower Kufrah River paleodrainage system. New analysis of SRTM-derived topography combined with Synthetic Aperture Radar images from the Japanese PALSAR orbital sensor allowed the mapping of seven main paleochannels located west of the Kufrah River, each of which is likely to have formed a tributary that supplied water and sediment to the main paleodrainage system. The northernmost four paleochannels probably originated from the Al Haruj relief, a Pliocene alkaline basaltic intra-continental volcanic field, and potentially connected to the Wadi Sahabi paleochannel. The remaining three paleochannels are in the more southerly location of the Sarir Calanscio, northeast of the Tibesti mountains, and barely present a topographic signature in SRTM data. They end in the dunes of the Calanscio Sand Sea, forming alluvial fans. The most southern paleochannel, known as Wadi Behar Belama, was previously mapped by Pachur et al. (1996) using LANDSAT-TM images, and was interpreted by Osborne et al. (2008) as representing part of an uninterrupted sediment pathway from the Tibesti mountains to the Mediterranean Sea. Processing of SRTM topographic data revealed local depressions which allow to connect the seven paleochannels and possibly the terminal alluvial fan of the Kufrah River to the Wadi Sahabi paleochannel, through a 400 km-long, south-north oriented, paleocorridor. These new findings support our previous hypothesis that proposed a connection between the lower Kufrah River in the region of the Sarir Dalmah and the Wadi Sahabi paleochannel, which connected to the Mediterranean Sea. Including the newly mapped paleochannels, the Kufrah River paleowatershed, at its maximum extent, would have covered more than 400000 km², representing close to a quarter of the surface area of Libya.

59
60
61
62
63
64
65

Keywords: Kufrah River, paleodrainage system, Wadi Sahabi, Libya, PALSAR, SRTM.

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56

Résumé : Paillou et al. (2009) ont cartographié un ancien réseau hydrographique de 900 km de long en Libye orientale, la rivière de Kufrah, qui a potentiellement relié le bassin de Kufrah à la mer Méditerranée au Miocène. Nous étudions ici la possible connexion entre la partie terminale de la rivière de Kufrah et la mer Méditerranée, via le paléochenal Wadi Sahabi, qui a potentiellement joué le rôle de partie terminale du système hydrographique de Kufrah. L'analyse de données topographiques SRTM, combinées à des images radar issues du capteur orbital japonais PALSAR, a permis de cartographier sept paléochenaux principaux à l'ouest de la rivière de Kufrah, chacun ayant potentiellement joué le rôle d'affluent pour le réseau hydrographique principal. Les quatre paléochenaux les plus au nord sont probablement originaires des reliefs du Al Haruj, un champ volcanique basaltique intra-continentale du Pliocène, et se connectent au paléochenal Wadi Sahabi. Les trois paléochenaux les plus au sud sont situés dans le Sarir Calanscio, au nord-est du Tibesti, et présentent une signature topographique difficilement détectable dans les données SRTM. Ils se terminent dans les dunes de la mer de sable de Calanscio, formant des cônes alluviaux. Le paléochenal le plus méridional, connu sous le nom de Wadi Behar Belama, a été précédemment cartographié par Pachur et al. (1996) à partir d'images LANDSAT-TM. Il a été considéré par Osborne et al. (2008) comme le segment d'un couloir sédimentaire continu qui reliait les montagnes du Tibesti à la mer Méditerranée. Le traitement des données topographiques SRTM a permis de mettre en évidence des dépressions locales qui permettent de relier les sept paléochenaux, et potentiellement le cône alluvial terminal de la rivière de Kufrah, au paléochenal Wadi Sahabi, et ce via un paléocorridor de 400 km de long, orienté sud-nord. Ces résultats confortent notre hypothèse d'une connexion entre la partie terminale de la rivière de Kufrah dans le Sarir Dalmah et la mer Méditerranée, via le paléochenal Wadi Sahabi. En tenant compte des sept nouveaux paléochenaux cartographiés, le bassin versant de la rivière de Kufrah aurait couvert plus de 400000 km², représentant presque un quart de la surface actuelle de la Libye.

57
58
59
60
61
62
63
64
65

Mots clés : Rivière de Kufrah, paléo-hydrographie, Wadi Sahabi, Libye, PALSAR, SRTM.

1. Introduction

1 While the central Sahara is now hyper-arid, extensive paleodrainage systems originating in
2 the Tibesti mountains used to flow northward to the Mediterranean Sea and southward to the
3
4 Chad Basin during wetter periods (Petit-Maire, 1998; Maley, 2010). Evidence of such
5
6 paleodrainage systems have been detected using various remote sensing imagery (Griffin,
7
8 2006; Pachur and Altmann, 2006; Drake et al., 2008), particularly by orbital imaging radar,
9
10 which allows the detection of paleochannels even when masked by Quaternary aeolian
11
12 deposits (McCauley et al., 1982; Abdelsalam and Stern, 1996; Robinson et al., 2006; Paillou et
13
14 al., 2010). Griffin (2002) proposed a vast paleodrainage system, the “Sahabi River system”
15
16 (see Figure 1), which would have flowed from the Messinian (late Miocene) Lake Chad,
17
18 eroded the east Tibesti valley, and have ended in a well preserved channel near the coast of
19
20 the Gulf of Sirt, the Wadi Sahabi paleochannel (Barr and Walker, 1973; Hallet, 2002; Swezey,
21
22 2009). The proposed path for the Sahabi River is poorly defined, however, and is based
23
24 mainly on the interpretation of low resolution topographic maps and on hypothetical
25
26 fragments of river channels detected in LANDSAT-TM images (Griffin, 2006; Griffin 2011).
27
28 Drake et al. (2008) also proposed that the Gulf of Sirt was fed by large river systems,
29
30 originating in northern and eastern Tibesti, through deep canyons that drained much of Libya
31
32 during the late Miocene. Although it was not mapped precisely, Drake et al. (2008) supported
33
34 the idea of the Sahabi River system that connected northern Chad and southeast Libya to the
35
36 Mediterranean Sea during humid periods in the Messinian. They also proposed the hypothesis
37
38 that this paleodrainage system was later captured by a more easterly one, the “River Al
39
40 Kufrah”, which was activated by tectonic subsidence in the Kufrah Basin during the Pliocene,
41
42 and linked to the Mediterranean Sea through the Wadi Sahabi paleochannel in the Sirt Basin.
43
44 The geographical mapping of this Kufrah River was initiated by Robinson et al. (2006) and
45
46 completed by Paillou et al. (2009) (see Figure 1). Ghoneim et al. (2012) confirmed the
47
48 geographical extent of the Kufrah River and proposed that its southwestern branch may have
49
50 served as an outlet from the Megalake Chad to the Mediterranean Sea during humid phases of
51
52 the Neogene. However, the hypothesis of a Sahabi River and/or Kufrah River system that
53
54
55
56
57
58
59
60
61
62
63
64
65

1 would provide a path for the Megalake Chad to discharge into the Mediterranean Sea is still a
2 matter of debate since there is yet no clear evidence of a connection between the Chad and
3 Kufrah basins, and other authors propose an evacuation of the overflow waters to the southern
4 Niger (Maley, 2004; Leblanc et al., 2006; Schuster et al., 2009; Maley, 2010). The results we
5 present in this study are not to be interpreted as supporting nor refuting the hypothesis of a
6 continuous connection between the Chad basin and the Mediterranean Sea, since we rather
7 focus here on the relationships between the lower Kufrah River and northern Tibesti and Al
8 Haruj reliefs, and propose possible connections to the Mediterranean Sea based on the
9 mapping of actual paleochannels and on the analysis of topography. In a same quantitative
10 mapping approach, Pachur and colleagues have also proposed and partially mapped an
11 extensive paleodrainage system that could have connected the Tibesti mountains to the Sirt
12 Basin during the Holocene (Pachur, 1980; Pachur, 1996; Pachur and Hoelzmann, 2000). This
13 paleodrainage system was divided into two main parts: One western part that was sourced in
14 the northern Tibesti mountains and then flowed to the north through the Behar Belama
15 paleochannel, and an eastern part that possibly originated in the eastern Tibesti mountains,
16 flowed into the Kufrah Basin, followed the present-day Wadi Blittah and ended as an inland
17 delta in the Sarir Dalmah during the Holocene (Pachur, 1996; Pachur and Altmann, 2006).

18 We present here the mapping of possible new tributaries of the Kufrah River, with their
19 potential flow path to the Mediterranean Sea, and discuss the implications for establishing an
20 extension of the watershed of this major Libyan paleodrainage system. Our results confirm the
21 previous work by Pachur and colleagues and we propose in addition some potential water flow
22 paths to the terminal Wadi Sahabi paleochannel. Although we have no new geochronological
23 or geochemical data, if it can be determined that these tributaries were active at intervals
24 during the late Pleistocene, then our findings would lend strong support to the previous
25 hypothesis of Rohling et al. (2002) and Osborne et al. (2008), who have proposed a continuous
26 “humid corridor” between southeastern Libya and the Mediterranean Sea at around 120 ka.

2. Mapping of the Kufrah River using Orbital Imaging Radar

Using data from the PALSAR L-band orbital radar of the Japanese ALOS satellite (Rosenqvist et al., 2007), we are conducting continental-scale mapping of several arid regions on Earth, with an initial objective to evaluate potentials of low frequency radar for planetary exploration (Paillou et al., 2006). A mosaic of the whole Sahara has been built from PALSAR strips, allowing an efficient detection of sub-surface geological features, particularly craters and paleodrainage networks (Paillou et al. 2003; Paillou et al. 2004; Paillou et al., 2006; Paillou et al., 2009; Paillou et al., 2010), because radar can probe through the first few meters of superficial eolian deposits (McCauley et al., 1982; Schaber et al., 1986; Paillou et al., 2003; Baghdadi et al., 2005; Grandjean et al., 2006). We thus precisely mapped a 900 km-long paleodrainage system in eastern Libya, termed the Kufrah River (see Figure 1), which could have linked the Kufrah Basin to the Mediterranean coast through Wadi Sahabi paleochannel in the Sirt Basin, possibly as far back as the middle Miocene (Paillou et al., 2009). The headwaters of this paleodrainage system are mainly in southern Libya, with detected tributaries arising in three main areas: El Fayoud and El Akdamin hamadas in northeastern Tibesti, northern Uweinat close to the Sudanese border, and the western Gilf Kebir and Abu Ras plateaux on the Egyptian border (see Figure 1). The Tibesti and Uweinat tributaries, more than 350 km-long, flowed in wide paleovalleys which join in the present-day Kufrah oasis. These paleovalleys had previously been detected by Robinson et al. (2006) using RADARSAT-1 C-band orbital radar. About 80 km north-east of the Kufrah oasis, a shorter (200 km-long) tributary joins from the Gilf Kebir and Abu Ras plateaux. From the Kufrah oasis, the main paleochannel becomes narrower (less than 1 km) and clearly incises the sandstone bedrock. It follows the present-day Wadi Blittah to the northern Jebel Dalmah over a distance of about 230 km. Farther north, in the Sarir Dalmah, the Kufrah River then disperses as a network of small, shallow paleochannels across the surface of a broad alluvial fan that covers more than 15000 km² (see Figure 1). It is not possible to follow the paleodrainage course to the north, because the large sand dunes of the Calanscio Sand Sea preclude radar mapping of the sub-surface. However, about 300 km away to the north-west and emerging from beneath the Calanscio

1 Sand Sea, lies the major, 2 to 4 km-wide, alluvium-filled Wadi Sahabi paleochannel that
2 incised more than 300 meters into Miocene carbonate strata (Barr and Walker, 1973). In a
3 previous study (Paillou et al., 2009), we proposed that the sand sea could hide an ancient
4 pathway between the Sarir Dalmah alluvial fan and the Wadi Sahabi paleochannel (cf. “north
5 path” marked by the red dotted line in Figure 1).
6
7
8
9
10

11 **3. New Tributaries from PALSAR and SRTM Data**

12
13
14
15
16 In addition to our PALSAR mosaic of the Sahara, we have used topographic information
17 derived from SRTM data. The Shuttle Radar Topography Mission consisted of an
18 interferometric radar system that flew on-board the Space Shuttle Endeavour during an 11-day
19 mission in February 2000 (Farr et al., 2007). It produced a high-quality global Digital
20 Elevation Model at a resolution of 3 arc-second (about 90 m), covering all land between
21 latitudes 56°S and 60°N. SRTM data are organized in 1° x 1° cells, and have an absolute
22 vertical height accuracy better than 15 meters. The HydroSHEDS (Hydrological data based on
23 SHuttle Elevation Derivatives at multiple Scales) data set was computed from SRTM
24 topography data by the USGS, and provides hydrographic information such as river networks,
25 watershed boundaries and drainage directions (Lehner et al., 2008). Due to errors and voids in
26 SRTM data, river network products are susceptible to various errors, particularly in very low
27 relief regions. This is unfortunately the case in our region of interest, where low relief
28 combines with numerous voids in SRTM coverage (mainly due to the radar wave penetration
29 and attenuation in sand dunes), and so does not enable the reliable generation of river networks
30 in HydroSHEDS. However, the computing of watershed boundaries is less sensitive to such
31 problems, and watershed boundaries available in HydroSHEDS data can be useful for
32 predicting areas to prospect for paleochannels using radar imagery. In particular, it appears that
33 a large drainage basin exists west of the lower Kufrah River, arising in the Sarir Tibesti and
34 flowing north in the Sarir Calanscio (see Figure 1): A paleodrainage contribution from the
35 western topography (Al Haruj area and northern Tibesti mountains) should then be expected.
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60
61
62
63
64
65

1 Combining topography derived from the SRTM data and the sub-surface imaging
2 capabilities of the PALSAR sensor, we mapped seven new paleochannels that are likely to
3 have contributed water and sediment to the Kufrah River paleodrainage system from the west
4 (see Figure 2). The northernmost four paleochannels probably originated from the Al Haruj
5 relief, a Pliocene basaltic intra-continental volcanic field, and potentially connected to the
6 northern Wadi Sahabi paleochannel. The remaining three paleochannels, which include the
7 Wadi Behar Belama paleochannel previously mapped by Pachur and Altmann (2006), are
8 found in the more southerly region of the Sarir Calanscio, located northeast of the Tibesti
9 mountains. They end in the Calanscio Sand Sea, forming alluvial fans that are oriented and
10 slope in the direction of the hypothesized but as yet unmapped 300 km-long “north path” of the
11 Kufrah River (Figure 1).
12
13
14
15
16
17
18
19
20
21
22
23
24
25

26 ***3.1. The Northern Paleochannels***

27
28 We mapped four main northern paleochannels (see Figure 2), which are narrow, single and
29 essentially straight and range from 110 to 190 km-long. They are associated with clear
30 topographic depressions in SRTM data, ranging between 10 and 30 meters in depth, and the
31 corresponding PALSAR images show well defined and narrow (less than 400 m) dark
32 channels. The paleochannels all appear to start on a limestone plateau that borders the
33 volcanic relief of Al Haruj, a young (~6 to 0.5 Ma) alkaline basaltic intra-continental volcanic
34 field (Ade-Hall et al., 1974). Although we cannot map the paleochannels as far west as the Al
35 Haruj relief because of sand dunes, they are very likely to have originated there. All the
36 paleochannels would have flowed from the southwest to the northeast and appear to terminate
37 about 30 km to the south of a southerly tributary of the Wadi Sahabi paleochannel, as shown
38 on Figure 3. Although we have no indication about the age of these paleochannels, it is very
39 likely that their formation is related to the Pliocene volcanic events that led to the formation
40 of the Al Haruj relief (Farahat et al., 2006).
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60
61
62

63 ***3.2. The Southern Paleochannels***

64
65

We mapped three main southern paleochannels, which are also narrow (less than 200 m) and essentially straight, but longer (200 to 300 km) than the northern ones (see Figure 2). While they barely present a topographic signature in SRTM data, the PALSAR radar sensor clearly detects them. These paleochannels traverse the Sarir Calanscio plain, which descends northeastwards from the Tibesti mountains (see Figure 2). The three southern paleochannels appear to emerge from the dune fields that cover the Sarir Tibesti: Potentially, they could be extended further southwest towards the Tibesti mountains, but the PALSAR sensor is not able to detect and map them through the thick sand cover. The most southerly paleochannel was previously mapped under the name of Wadi Behar Belama (see Figure 2) by Pachur and Altmann (2006) using LANDSAT-TM images. As with the northernmost paleochannels, the southernmost ones follow a southwest to northeast direction. They terminate at the western margin of the Calanscio Sand Sea, dispersing as networks of small, shallow paleochannels across the surface of alluvial fans (see Figure 4), indicating a decrease in paleoflow competence. Coarse alluvial sand and gravel constituting the fan sediments increase the surface roughness and volume scattering effects, so that the southern paleochannel terminations appear as bright features in the radar images, as shown on Figure 4.

4. Topography Analysis and Discussion

The seven paleochannels that we have mapped are consistent with the hypothesis of an extensive Kufrah River paleodrainage system that connected the northern Tibesti mountains to the Sirt Basin during the late Cenozoic as previously proposed by Pachur et al. (Pachur, 1980; Pachur, 1996; Pachur and Hoelzmann, 2000; Pachur and Altmann, 2006). Using LANDSAT-TM imagery, Pachur and colleagues detected remnants of a paleodrainage system formed by the junction of former great wadis originating in the northern Tibesti mountains, that crossed the Sarir Calanscio in a northeast direction. A main representative of this system is Wadi Behar Belama which contains acid volcanic pebbles originating from the Tibesti. Pachur and colleagues also mapped the upper reaches of a paleoriver system sourced in the Al

1 Haruj volcanic relief, that “probably joined drainage systems that flowed into the
2 Mediterranean” (Pachur, 1996), and suggested that this system is likely to have been active
3 during the early Holocene. The new paleochannels that we have mapped are also consistent
4 with the location of the “humid corridor” proposed by Rohling et al. (2002) and Osborne et al.
5 (2008), that was connecting southeastern Libya to the Mediterranean Sea at around 120 ka.
6 Osborne et al. (2008) performed a Nd isotopic characterization of Quaternary sediments
7 sampled in Wadi Behar Belama and used the findings to support the interpretation of an
8 uninterrupted sediment transport pathway from the Tibesti mountains to the Mediterranean
9 Sea. The alluvial fans at the termini of Wadi Behar Belama and of the two other southern
10 paleochannels (see Figure 2), might then be transitional features produced during drier
11 periods, when the paleoflow competence decreased and sediment throughput could not be
12 maintained. These paleochannels have certainly contributed to sediment supply of the
13 Calanscio Sand Sea (El Baz et al., 2000; Ghoneim et al., 2012) and so may have also
14 contributed to burial of the through-going “north path” of the main Kufrah River that we
15 hypothesized (Paillou et al., 2009).

16 While full resolution SRTM topography contains very detailed features, particularly
17 dunes, that preclude detailed mapping of the main paleochannels, one can consider the low-
18 frequency topographic information in order to better infer paleodrainage directions. Craddock
19 et al. (2010) applied this approach to reconstruct a buried fluvial landscape beneath the
20 Simpson Desert eolian dune field in central Australia: While many dunes screen the main
21 drainage directions in full resolution SRTM data, sampling of the low-frequency topography
22 in the topographic lows between the dunes revealed slope trends that trace remnants of
23 paleodrainages. We applied a similar approach here by low-pass filtering the SRTM data in
24 the frequency domain, after a Fourier transform, and by using interpolation techniques to fill
25 SRTM voids: Figure 5a shows the full resolution SRTM topography with voids filled by
26 interpolation, and Figure 5b shows the low-frequency topography obtained after low-pass
27 filtering. Applying a simple threshold to the low-frequency topography allows to directly map
28 the local depressions, represented as light blue areas in Figure 5c. There is no clear evidence

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60
61
62
63
64
65

of a topographic depression that would extend the alluvial fan terminating the Kufrah River to the northwest: The hypothetical “north path” connecting the Sarir Dalmah to the Wadi Sahabi paleochannel as indicated on Figure 1 cannot be observed on present-day topography (but still could be buried under the dunes of the Calanscio Sand Sea). However, a local depression (path #1 in Figure 5c) seems to connect the Kufrah River alluvial fan to a western corridor (path #2 in Figure 5c). The latter runs from the northern Tibesti, through the Wadi Behar Belama in the Sarir Calanscio, and ends south of the Wadi Sahabi paleochannel (see Figure 5c). This large paleocorridor, 10 to 20 km wide and about 400 km long, is close to the terminating path of the lower Sahabi River as proposed by Griffin and Drake (see Figure 1) but again, as stated in the introduction of this study, this cannot be considered as a proof for a continuous river system connecting the Chad basin to the Mediterranean Sea. The observed paleocorridor is more likely to have been fed by the northern Tibesti and Al Haruj reliefs, being an ancient pathway between the paleochannels described in the previous section and the northern Wadi Sahabi paleochannel (see Figure 5c). This paleocorridor, as a depression adjacent to the margins of the aggraded Sarir Dalmah alluvial fan, was possibly re-enforced by local tectonics, since its orientation is comparable to the orientation of the main faults in this area (Ahlbrandt, 2001). In particular, one can observe in Figure 3 that the course of two of the northernmost paleochannels shows a change in the paleoflow direction of about 90° when approaching the paleocorridor from the southwest, suggesting a possible tectonic control by the horst and graben structures in the region (Ahlbrandt, 2001).

The newly mapped paleochannels and the analysis of the local topography thus strongly support the hypotheses proposed by Pachur (1980, 1996), Rohling et al. (2002) and Osborne et al. (2008) of formerly continuous connections between drainage basins in northern Tibesti and the Mediterranean Sea through the Wadi Sahabi paleochannel. Our findings also support our previous hypothesis that the Sarir Dalmah alluvial fan, located at the terminus of the upper part of the Kufrah River system, could have connected to the Wadi Sahabi paleochannel in the north (Paillou et al., 2009). This connection is however not yet obvious, and the Kufrah River is also likely to have terminated as in inland delta in the Sarir Dalmah

(Paillou et al., 2009), possibly feeding the northern Al Jaghbub paleolake as proposed by Ghoneim et al. (2012). This re-emphasizes the importance of undertaking some exploratory field work in the Calanscio Sand Sea, using geophysical prospecting techniques such as Ground Penetrating Radar, to detect and map possible paleochannels buried under the sand dunes. The Calanscio Sand Sea is a key area for understanding the history of paleodrainage systems of eastern Libya, and near surface geophysical prospecting could provide new data to help answer outstanding questions regarding the connection of the paleochannels in this area. Exploratory field work is also needed in order to collect samples for geochronological and geochemical analyses that would help determine the ages and histories of the various paleochannels that we have mapped so far: The many tributaries of the Kufrah River system may not necessarily have the same antiquity, some of them possibly being younger impositions on an older paleodrainage network (in particular, the northern paleochannels related to the young volcanic relief of Al Haruj are certainly more recent than the southern ones).

With the plausible assumption that all the seven newly mapped paleochannels at one time were forming part of the Kufrah River paleodrainage system, we can now define a major paleodrainage system in eastern Libya, which at its maximum extent would have drained an area of more than 400000 km² between the Tibesti, Al Haruj and Gilf Kebir massifs and probably connected to the Mediterranean Sea through the Wadi Sahabi paleochannel in the Sirt Basin. The whole system is likely to have been active at intervals during the late Cenozoic, possibly discharging at that time a comparable amount of water as does the present-day Nile into the Mediterranean Sea. The Kufrah River system is then clearly a major paleohydrological feature to take into account when studying the past environments and climates of northern Africa from the middle Miocene to the Holocene. It also represents a likely corridor for fauna and human dispersal in the eastern Sahara, and thus indicates locations where further paleontological, paleo-anthropological and archeological field exploration should be conducted.

Acknowledgments

The authors would like to thank JAXA (Japan Aerospace Exploration Agency) for providing PALSAR data in the framework of the Kyoto & Carbon Initiative. We also thank P. deMenocal, E. Rohling, J. Maley, Ph. Düringer, Gh. de Marsily and three anonymous reviewers for their very constructive comments on earlier versions of this manuscript. This work was financially supported by the French space agency CNES (Centre National d'Etudes Spatiales).

References

- Abdelsalam M. G. and R. J. Stern, "Mapping precambrian structures in the Sahara Desert with SIR-C/X-SAR radar: The neoproterozoic Keraf suture, NE Sudan," *J. Geophys. Res.*, vol. 101, no. E10, pp. 23063-23076, 1996.
- Ade-Hall F. M., P. H. Reynolds, P. Dagley, A. G. Musset, T. B. Hubbard, E. Klitzsch, "Geophysical studies of North African Cenozoic volcanic areas A1-Haruj Assuad, Libya," *Canadian Journal of Earth Science*, vol. 11, pp. 998-1006, 1974.
- Ahlbrandt T. S., "The Sirte Basin province of Libya – Sirte-Zelten Total Petroleum system", *U.S. Geological Survey Bulletin 2202-F*, 29 p., 2001.
- Baghdadi N., G. Grandjean, D. Lahondère, Ph. Paillou, Y. Lasne, "Apport de l'imagerie satellitaire radar pour l'exploration géologique en zone aride," *C.R. Geoscience*, vol. 337, pp. 719-728, 2005.
- Barr F. T. and B. R. Walker, "Late Tertiary channel system in northern Libya and its implications on Mediterranean sea level changes," In: W. B. F. Ryan et al.. (Eds.), *Init. Rep. DSDP 13*, pp. 1244-1251, 1973.
- Craddock R. A., M. F. Hutchinson, J. A. Stein, "Topographic data reveal a buried fluvial landscape in the Simpson Desert, Australia," *Australian Journal of Earth Sciences*, vol. 57, pp. 163-171, 2010.
- Drake N. A., A. S. El-Hawat, P. Turner, S. J. Armitage, M. J. Salem, K. H. White, S. McLaren, "Palaeohydrology of the Fazzan Basin and surrounding regions: The last 7 million years," *Palaeogeography, Palaeoclimatology, Palaeoecology*, vol. 263, pp. 131-145, 2008.

- 1 El Baz F., Mainguet M., Robinson C., "Fluvio-aeolian dynamics in the north-eastern Sahara: The
2 relationship between fluvial/aeolian systems and ground water concentration," *J. of Arid*
3 *Environments*, vol. 44, pp. 173-183, 2000.
- 4 Farahat E. S., M. S. Abdel Ghani, A. S. Aboazom, A. M. H. Asran, "Mineral chemistry of Al Haruj
5 low-volcanicity rift basalts, Libya: Implications for petrogenetic and geotectonic evolution,"
6 *Journal of African Earth Sciences*, vol. 45, pp. 198-212, 2006.
- 7
8
9
10 Farr T. G. et al., "The Shuttle Radar Topography Mission," *Rev. Geophys*, vol. 45, RG2004,
11 doi:10.1029/2005RG000183, 2007.
- 12
13
14 Ghoneim E., M. Benedetti, F. El Baz, "An integrated remote sensing and GIS analysis of the Kufrah
15 Paleoriver, Eastern Sahara," *Geomorphology*, vol. 139-140, pp. 242-247, 2012.
- 16
17
18 Grandjean G., Ph. Paillou, N. Bahgdadi, E. Heggy, T. August, Y. Lasne, "Surface and subsurface
19 structures mapping using low frequency radar: A synthesis of the Mauritanian and Egyptian
20 experiments," *Journal of African Earth Sciences*, vol. 44, no. 2, pp. 220-228, 2006.
- 21
22
23
24
25 Griffin D. L., "Aridity and humidity: two aspects of the late Miocene climate of North Africa and the
26 Mediterranean," *Palaeogeography, Palaeoclimatology, Palaeoecology*, vol. 182, pp. 65-91, 2002.
- 27
28
29
30 Griffin D. L., "The late Neogene Sahabi rivers of the Sahara and their climatic and environmental
31 implications for the Chad basin," *Journal of the Geological Society*, London, vol. 163, pp. 905-
32 921, 2006.
- 33
34
35
36
37 Griffin D. L., "The late Neogene Sahabi rivers of the Sahara and the hamadas of the eastern Libya-
38 Chad border area," *Palaeogeography, Palaeoclimatology, Palaeoecology*, vol. 309, pp. 176-185,
39 2011.
- 40
41
42
43 Hallet D., *Petroleum Geology of Libya*, Elsevier, Amsterdam, 503p., 2002.
- 44
45
46 Leblanc M., G. Favreau, J. Maley, et al., "Reconstruction of Megalake Chad using Shuttle Radar
47 Topographic Mission data," *Palaeogeography, Palaeoclimatology, Palaeoecology*, vol. 239, pp.
48 16-27, 2006.
- 49
50
51
52
53
54
55
56
57
58
59
60
61
62
63
64
65
66
67
68
69
70
71
72
73
74
75
76
77
78
79
80
81
82
83
84
85
86
87
88
89
90
91
92
93
94
95
96
97
98
99
100
101
102
103
104
105
106
107
108
109
110
111
112
113
114
115
116
117
118
119
120
121
122
123
124
125
126
127
128
129
130
131
132
133
134
135
136
137
138
139
140
141
142
143
144
145
146
147
148
149
150
151
152
153
154
155
156
157
158
159
160
161
162
163
164
165
166
167
168
169
170
171
172
173
174
175
176
177
178
179
180
181
182
183
184
185
186
187
188
189
190
191
192
193
194
195
196
197
198
199
200
201
202
203
204
205
206
207
208
209
210
211
212
213
214
215
216
217
218
219
220
221
222
223
224
225
226
227
228
229
230
231
232
233
234
235
236
237
238
239
240
241
242
243
244
245
246
247
248
249
250
251
252
253
254
255
256
257
258
259
260
261
262
263
264
265
266
267
268
269
270
271
272
273
274
275
276
277
278
279
280
281
282
283
284
285
286
287
288
289
290
291
292
293
294
295
296
297
298
299
300
301
302
303
304
305
306
307
308
309
310
311
312
313
314
315
316
317
318
319
320
321
322
323
324
325
326
327
328
329
330
331
332
333
334
335
336
337
338
339
340
341
342
343
344
345
346
347
348
349
350
351
352
353
354
355
356
357
358
359
360
361
362
363
364
365
366
367
368
369
370
371
372
373
374
375
376
377
378
379
380
381
382
383
384
385
386
387
388
389
390
391
392
393
394
395
396
397
398
399
400
401
402
403
404
405
406
407
408
409
410
411
412
413
414
415
416
417
418
419
420
421
422
423
424
425
426
427
428
429
430
431
432
433
434
435
436
437
438
439
440
441
442
443
444
445
446
447
448
449
450
451
452
453
454
455
456
457
458
459
460
461
462
463
464
465
466
467
468
469
470
471
472
473
474
475
476
477
478
479
480
481
482
483
484
485
486
487
488
489
490
491
492
493
494
495
496
497
498
499
500
501
502
503
504
505
506
507
508
509
510
511
512
513
514
515
516
517
518
519
520
521
522
523
524
525
526
527
528
529
530
531
532
533
534
535
536
537
538
539
540
541
542
543
544
545
546
547
548
549
550
551
552
553
554
555
556
557
558
559
560
561
562
563
564
565
566
567
568
569
570
571
572
573
574
575
576
577
578
579
580
581
582
583
584
585
586
587
588
589
590
591
592
593
594
595
596
597
598
599
600
601
602
603
604
605
606
607
608
609
610
611
612
613
614
615
616
617
618
619
620
621
622
623
624
625
626
627
628
629
630
631
632
633
634
635
636
637
638
639
640
641
642
643
644
645
646
647
648
649
650
651
652
653
654
655
656
657
658
659
660
661
662
663
664
665
666
667
668
669
670
671
672
673
674
675
676
677
678
679
680
681
682
683
684
685
686
687
688
689
690
691
692
693
694
695
696
697
698
699
700
701
702
703
704
705
706
707
708
709
710
711
712
713
714
715
716
717
718
719
720
721
722
723
724
725
726
727
728
729
730
731
732
733
734
735
736
737
738
739
740
741
742
743
744
745
746
747
748
749
750
751
752
753
754
755
756
757
758
759
760
761
762
763
764
765
766
767
768
769
770
771
772
773
774
775
776
777
778
779
780
781
782
783
784
785
786
787
788
789
790
791
792
793
794
795
796
797
798
799
800
801
802
803
804
805
806
807
808
809
810
811
812
813
814
815
816
817
818
819
820
821
822
823
824
825
826
827
828
829
830
831
832
833
834
835
836
837
838
839
840
841
842
843
844
845
846
847
848
849
850
851
852
853
854
855
856
857
858
859
860
861
862
863
864
865
866
867
868
869
870
871
872
873
874
875
876
877
878
879
880
881
882
883
884
885
886
887
888
889
890
891
892
893
894
895
896
897
898
899
900
901
902
903
904
905
906
907
908
909
910
911
912
913
914
915
916
917
918
919
920
921
922
923
924
925
926
927
928
929
930
931
932
933
934
935
936
937
938
939
940
941
942
943
944
945
946
947
948
949
950
951
952
953
954
955
956
957
958
959
960
961
962
963
964
965
966
967
968
969
970
971
972
973
974
975
976
977
978
979
980
981
982
983
984
985
986
987
988
989
990
991
992
993
994
995
996
997
998
999
1000

- Maley J., "Climate and palaeoenvironment evolution in north tropical Africa from the end of the Tertiary to the Upper Quaternary," *Palaeoecology of Africa*, vol. 30, pp. 227-278, 2010.
- 1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60
61
62
63
64
65
- McCauley J. F., G. G. Schaber, C. S. Breed, M. J. Grolier, C. V. Haynes, B. Issawi, C. Elachi, R. Blom., "Subsurface valleys and geoarchaeology of the eastern Sahara revealed by Shuttle Radar," *Science*, vol. 218, pp. 1004-1020, 1982.
- Osborne A. H., D. Vance, E. J. Rohling, N. Barton, M. Rogerson, N. Fello, "A humid corridor across the Sahara for the migration of early modern humans out of Africa 120,000 years ago," *Proc. Nat. Acad. Sci.*, vol. 105, pp. 16444-16447, 2008.
- Pachur H. J., "Climatic history in the Late Quaternary in southern Libya and the western Libyan desert," In: M. J. Salem and M. T. Busrewil (Eds.), *The Geology of Libya*, vol. 3, Academic Press, London, pp. 781-788, 1980.
- Pachur H. J., "The Geology of Syrte Basin - vol 1 - Reconstruction of paleodrainage systems in Syrte Basin and the area surrounding the Tibesti Mountains: Implications for the hydrological history of the region," *First symposium on the sedimentary basins of Libya*, Tripoli, Elsevier Eds., 1996.
- Pachur H. J. and P. Hoelzmann, "Late Quaternary palaeoecology and palaeoclimates of the eastern Sahara," *Journal of African Earth Sciences*, vol. 30, pp. 929-939, 2000.
- Pachur H. J. and N. Altmann, *Die Ostsahara im Spätquartär*, Springer Berlin Heidelberg New York, 662 p., 2006.
- Paillou Ph., G. Grandjean, N. Baghdadi, E. Heggy, Th. August-Bernex, J. Achache, "Sub-surface imaging in central-southern Egypt using low frequency radar: Bir Safsaf revisited," *IEEE Transactions on Geoscience and Remote Sensing*, vol. 41, no. 7, pp. 1672-1684, 2003.
- Paillou Ph., A. Rosenqvist, J.-M. Malézieux, B. Reynard, T. Farr, E. Heggy, "Discovery of a double impact crater in Libya: the astrobleme of Arkenu," *C.R. Geoscience*, vol. 335, pp. 1059-1069, 2003.
- Paillou Ph., A. El Barkooky, A. Barakat, J.-M. Malézieux, B. Reynard, J. Dejax, E. Heggy, "Discovery of the largest crater field on Earth in the Gilf Kebir region, Egypt," *C.R. Geoscience*, vol. 336, pp. 1491-1500, 2004.
- Paillou Ph., Y. Lasne, E. Heggy, J.-M. Malézieux, "A study of P-band SAR applicability and performance for Mars exploration: Imaging subsurface geology and detecting shallow moisture," *Journal of Geophysical Research*, vol. 111, E06S11, 2006.

- Paillou Ph., B. Reynard, J.-M. Malézieux, J. Dejax, E. Heggy, P. Rochette, W. U. Reimold, P. Michel, D. Baratoux, Ph. Razin, J.-P. Colin, "An extended field of crater-shaped structures in the Gilf Kebir region – Egypt: Observations and hypotheses about their origin," *Journal of African Earth Sciences*, vol. 46, pp. 281-299, 2006.
- Paillou Ph., M. Schuster, S. Tooth, T. Farr, A. Rosenqvist, S. Lopez, J.-M. Malézieux, "Mapping of a major paleodrainage system in Eastern Libya using orbital imaging Radar: The Kufrah River," *Earth and Planetary Science Letters*, vol. 277, pp. 327-333, doi: 10.1016/j.epsl.2008.10.029, 2009.
- Paillou Ph., S. Lopez, T. Farr, A. Rosenqvist, "Mapping Subsurface Geology in Sahara using L-band SAR: First Results from the ALOS/PALSAR Imaging Radar," *IEEE J. of Selected Topics in Earth Observations and Remote Sensing*, vol. 3, no 4, pp. 632-636, 2010.
- Petit-Maire N., "Climatic change and cultural change in the Sahara over the last 130 ka," In: *Il sistema Uomo-Ambiente tra passato e presente*, Albore Livadie C., Ortolani F. Eds., Centro universitario europeo per i beni culturali, pp. 311-316, 1998.
- Robinson C. A., F. El-Baz, T. S. M. Al-Saud, S. B. Jeon, "Use of radar data to delineate palaeodrainage leading to the Kufra oasis in the eastern Sahara," *Journal of African Earth Sciences*, vol. 44, pp. 229-240, 2006.
- Rohling E. J., T. R. Cane, S. Cooke, M. Sprovieri, I. Bouloubassi, K. C. Emeis, R. Schiebel, D. Kroon, F. J. Jorissen, A. Lorre, A. E. S. Kemp, "African monsoon variability during the previous interglacial maximum," *Earth and Planetary Science Letters*, vol. 202, pp. 61-75, 2002.
- Rosenqvist A., M. Shimada, N. Ito, M. Watanabe, "ALOS PALSAR: A pathfinder mission for global-scale monitoring of the environment," *IEEE Transactions on Geoscience and Remote Sensing*, vol. 45, no. 11, pp. 3307-3316, 2007.
- Schaber G. G., J. F. McCauley, C. S. Breed, and G. R. Olhoeft, "Shuttle Imaging Radar: Physical controls on signal penetration and subsurface scattering in the Eastern Sahara," *IEEE Transactions on Geoscience and Remote Sensing*, vol. GE-24, no. 4, pp. 603-623, 1986.
- Schuster M., Ph. Durringer, J.F. Ghienne, et al., "Chad Basin: Paleoenvironments of the Sahara since the Late Miocene," *C.R. Geoscience*, vol. 341, pp. 603-611, 2009.
- Swezey C. S., "Cenozoic stratigraphy of the Sahara, Northern Africa," *Journal of African Earth Sciences*, vol. 53, pp. 89-121, 2009.

Figure Captions

1
2 **Figure 1:** The Kufrah River paleodrainage system and the Wadi Sahabi paleochannel (in
3 blue) mapped onto a LANDSAT-TM mosaic. The hypothetical “north path” proposed by
4 Paillou et al. (2009) and “Sahabi River” proposed by Griffin (2002) and Drake (2008) are
5 indicated as dashed red lines.
6
7
8
9

10
11
12
13 **Figure 2:** The seven paleochannels, four northern ones and three southern ones – including
14 Wadi Behar Belama – which have been mapped using SRTM and PALSAR data. Background
15 is SRTM topography with interpolation to fill voids.
16
17
18
19
20
21
22

23 **Figure 3:** Northern limit of the northernmost paleochannels (thin blue) which are likely to
24 have connected to a southern tributary of the Wadi Sahabi paleochannel (thick blue), located
25 30 km to the north. Background is interpolated SRTM topography.
26
27
28
29
30
31
32

33 **Figure 4:** PALSAR image of the northern limit of the southern paleochannels. They terminate
34 at the western margin of the dunes of the Calanscio Sand Sea, dispersing as a network of
35 small, shallow paleochannels across the surface of radar-bright alluvial fans.
36
37
38
39
40
41
42

43 **Figure 5: a.** The Kufrah River, the Wadi Sahabi paleochannel and the seven newly mapped
44 paleochannels on top of the full resolution SRTM topography. **b.** The same features shown
45 over the low-frequency topography obtained after filtering in the Fourier domain. **c.** Local
46 depressions (light blue areas) extracted from the low-frequency topography. “Path #1” shows
47 a narrow valley which could connect the Sarir Dalmah alluvial fan to a paleocorridor “Path
48 #2”. This paleocorridor, 10 to 20 km wide and about 400 km long, runs from the northern
49 Tibesti, through the Wadi Behar Belama in the Sarir Calanscio, and ends south of the Wadi
50 Sahabi paleochannel.
51
52
53
54
55
56
57
58
59
60
61
62
63
64
65

Légende des figures

1
2 **Figure 1** : Le système paléo-hydrographique de la rivière de Kufrah et le paléochenal Wadi
3 Sahabi (en bleu) sur fond de mosaïque LANDSAT-TM. Le “north path” hypothétique
4 proposé par Paillou et al. (2009) et la “Sahabi River” proposée par Griffin (2002) et Drake
5 (2008) sont indiqués en traits pointillés rouges.
6
7
8
9

10
11
12
13 **Figure 2** : Les sept paléochenaux, quatre septentrionaux et trois méridionaux – incluant le
14 Wadi Behar Belama – qui ont été cartographiés à partir des données SRTM et PALSAR. Le
15 fond est la topographie SRTM, interpolée pour combler les trous.
16
17
18
19

20
21
22
23 **Figure 3** : Limite nord des paléochenaux septentrionaux (tracé bleu fin) qui semblent se
24 connecter à un affluent méridional du paléochenal Wadi Sahabi (tracé bleu épais), localisé 30
25 km plus au nord. Le fond est la topographie SRTM interpolée.
26
27
28
29

30
31
32
33 **Figure 4** : Image PALSAR de la limite nord des paléochenaux méridionaux. Ils se terminent
34 à la bordure occidentale des dunes de la mer de sable de Calanscio, se dispersant en un réseau
35 de petits chenaux, présentant une signature radar brillante, à la surface de cônes alluviaux.
36
37
38
39

40
41
42
43 **Figure 5** : **a.** La rivière de Kufrah, le paléochenal Wadi Sahabi et les sept nouveaux
44 paléochenaux sur fond de topographie SRTM. **b.** Les mêmes structures superposées à la
45 composante basse fréquence de la topographie, obtenue par filtrage dans le domaine de
46 Fourier. **c.** Dépressions locales (zones en bleu clair) extraites de la topographie basse
47 fréquence. Le “path #1” indique une vallée étroite qui connecte potentiellement le cône
48 alluvial du Sarir Dalmah à un paléocorridor “path #2”. Ce paléocorridor, large de 10 à 20 km
49 et long d'environ 400 km, débute dans la région nord du Tibesti, suit le Wadi Behar Belama
50 dans le Sarir Calanscio, et se termine au sud du paléochenal Wadi Sahabi.
51
52
53
54
55
56
57
58
59
60
61
62
63
64
65

Figure (pas au format Adobe Illustrator / Not Adobe Illustrator file)
[Click here to download high resolution image](#)

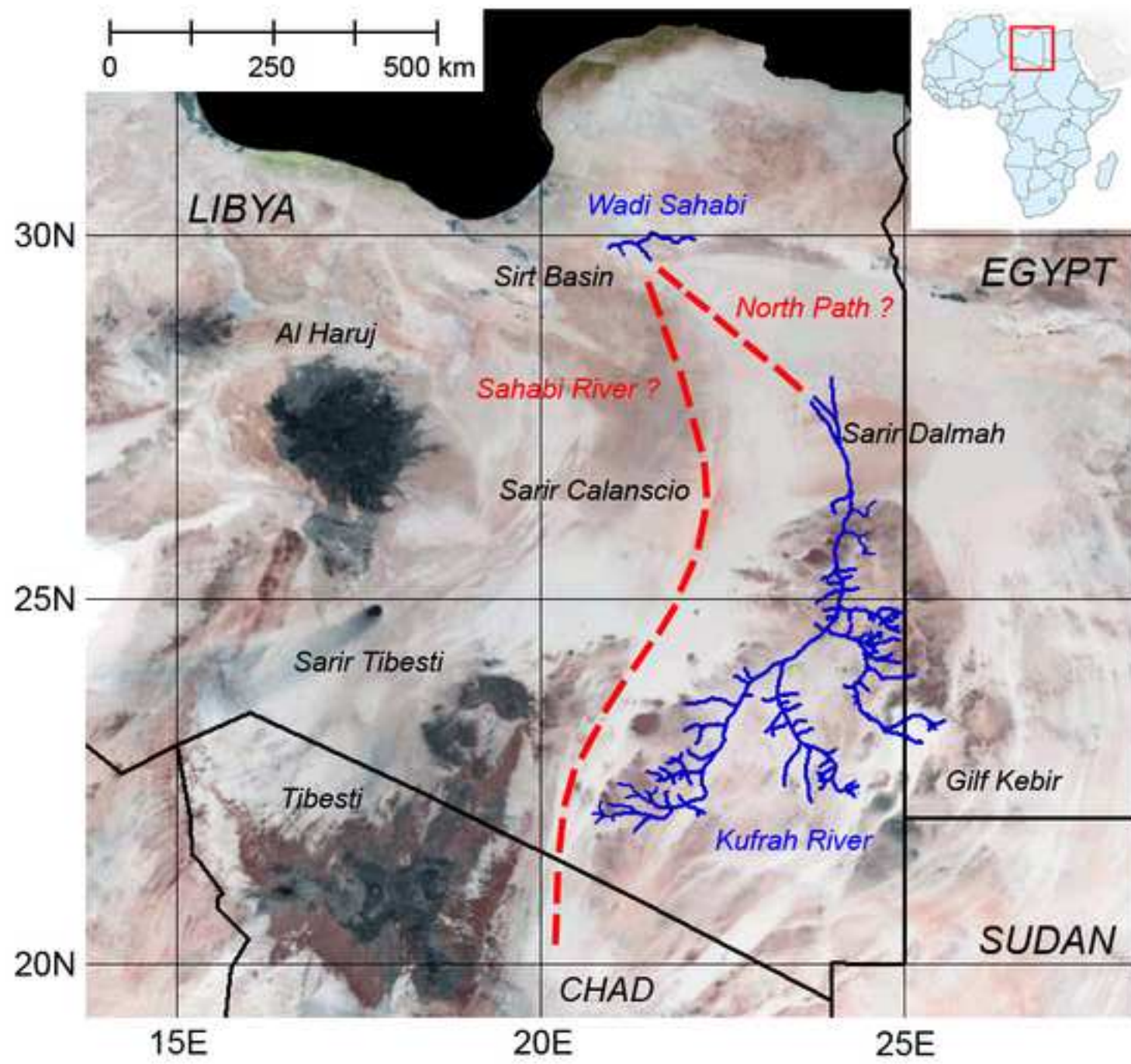


Figure (pas au format Adobe Illustrator / Not Adobe Illustrator file)
[Click here to download high resolution image](#)

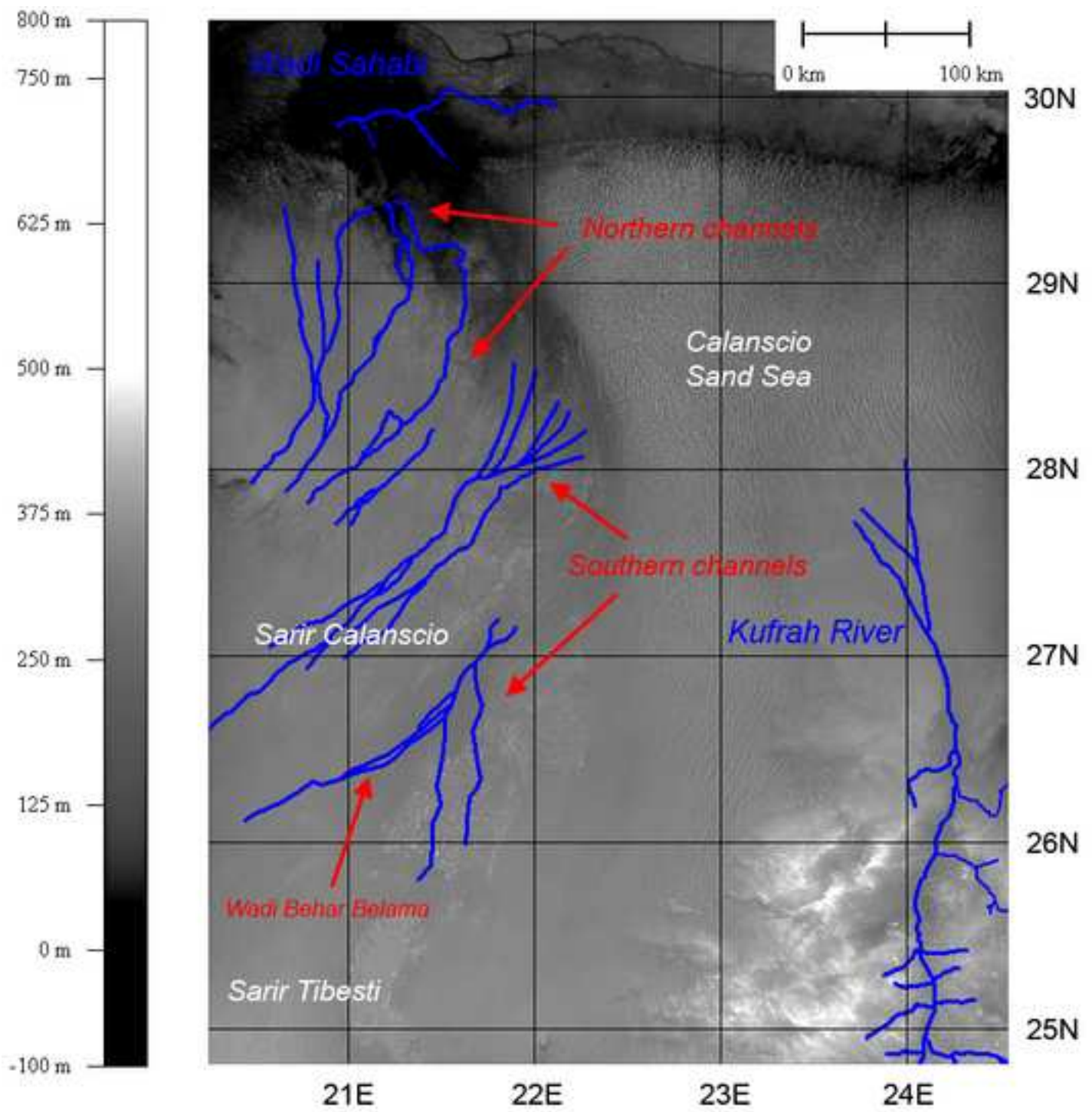


Figure (pas au format Adobe Illustrator / Not Adobe Illustrator file)
[Click here to download high resolution image](#)

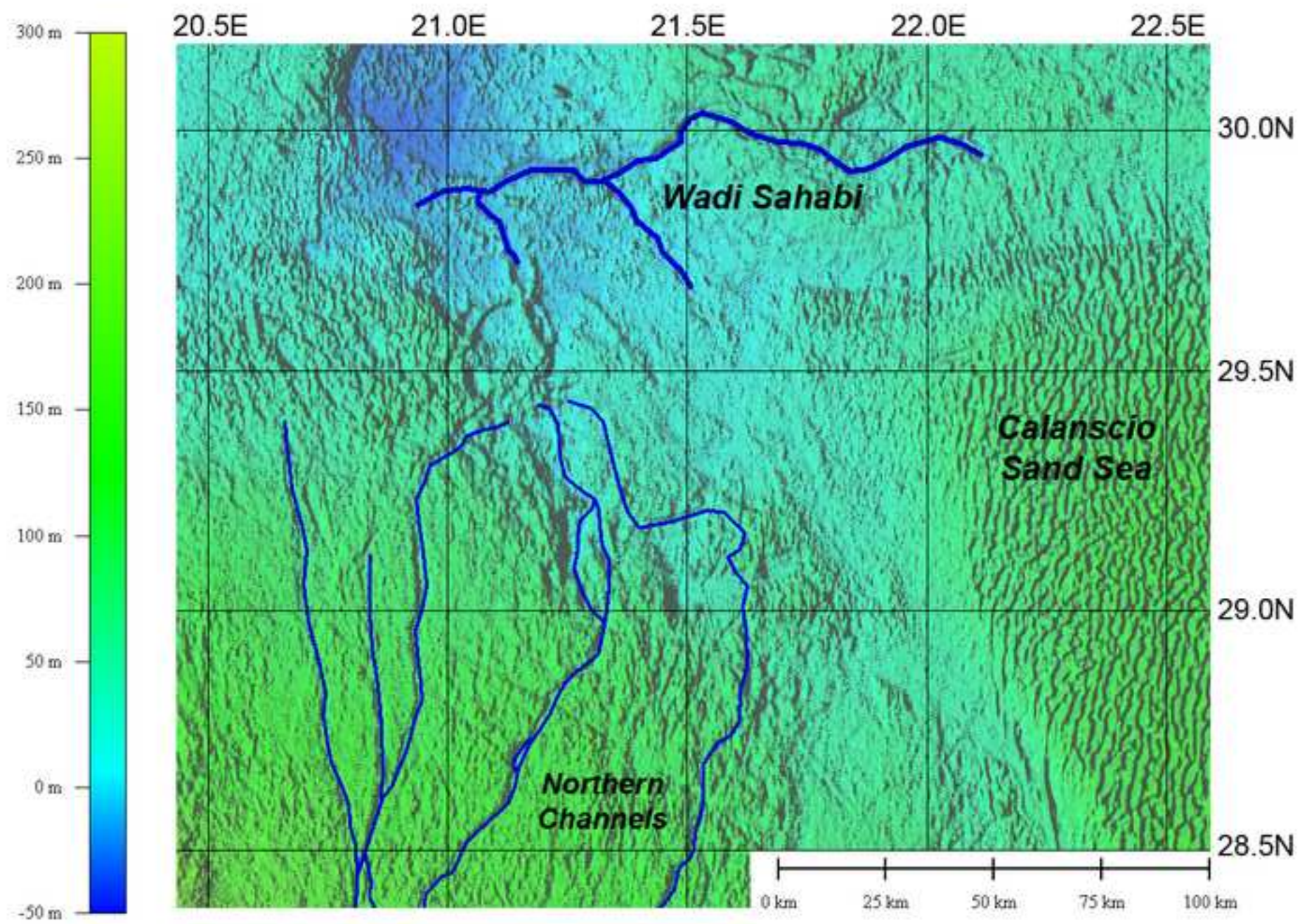


Figure (pas au format Adobe Illustrator / Not Adobe Illustrator file)
[Click here to download high resolution image](#)

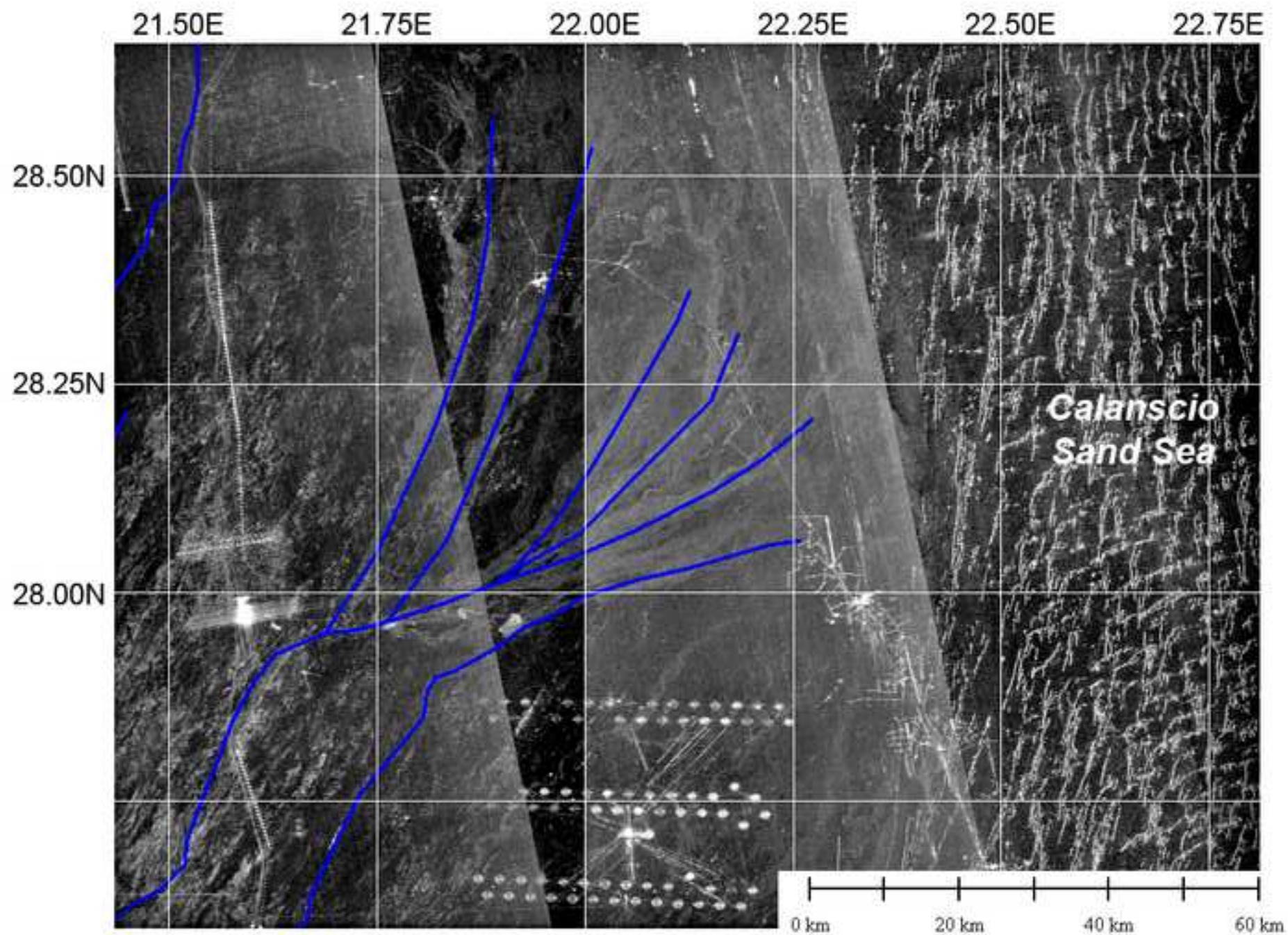


Figure (pas au format Adobe Illustrator / Not Adobe Illustrator file)
[Click here to download high resolution image](#)

